

**P.10.05.6***Acta Cryst.* (2005). A61, C377**Resolving Structure and Size of Amorphous Mineral Precipitates by PDF Analysis**

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Amorphous, or X-ray amorphous, is a common term used to describe mineral precipitates lacking long-range order. Earth materials, such as amorphous FeS, silica, and Fe(III)-hydroxide, are abundant in natural environments and play an important role in geochemical processes and the fate of contaminants. While use of this term persists in the Geosciences community, it is important to resolve the short- to medium-range structure and fundamental size of these solids to deepen our understanding of their role in geochemical processes. In this study, we report on the structure and size of so-called amorphous FeS, the first Fe-S phase formed in anoxic sediments which, over time, will convert to pyrite (FeS<sub>2</sub>). Amorphous FeS can also sequester metals and metalloids, such as cadmium and arsenic and react with contaminants, such as Cr<sup>6+</sup>. High-energy X-ray scattering data for freshly precipitated and aged FeS (7 days at 70°C) were collected at the Advanced Photon Source and analyzed using the PDF method. Results indicate that amorphous FeS has short- to medium-range order consistent with that of crystalline mackinawite. Further, the range of structural coherence (i.e., fundamental particle size) as can be obtained from PDF [1] is initially about 2 nm, but increases to 4.5 nm for the aged sample. Hence, the initial FeS precipitate is better described as a nanocrystalline material with mackinawite structure. This technique holds great promise for further study of nanocrystalline earth materials lacking long-range order.

[1] Page et al., *Chemical Physics Letters*, 2004, **393**, 385-388.

**Keywords:** pair distribution function, nanocrystals, particle size

**P.10.05.7***Acta Cryst.* (2005). A61, C377**Crystal structures of Rb<sub>2</sub>(C<sub>2</sub>O<sub>4</sub>)·H<sub>2</sub>O and Tl<sub>2</sub>(C<sub>2</sub>O<sub>4</sub>): Application of Valence Matching Rule**

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Crystal structures of Rb<sub>2</sub>(C<sub>2</sub>O<sub>4</sub>)·H<sub>2</sub>O (monoclinic, space group C2/c, *a* = 9.617(6)Å, *b* = 6.353(5)Å, *c* = 11.010(8)Å, β = 109.46(3)°, *V* = 634.2(8)Å<sup>3</sup>, *Z* = 4, *R*1 = 0.026, for 2646 reflections) and Tl<sub>2</sub>(C<sub>2</sub>O<sub>4</sub>) (triclinic, *P* 1̄, *a* = 6.623(4)Å, *b* = 6.674(3)Å, *c* = 5.854(4)Å, α = 90.031(35)°, β = 89.967(36)°, γ = 80.745(40)°, *V* = 255.3(3)Å<sup>3</sup>, *Z* = 2, *R*1 = 0.082, for 1499 reflections) were determined using an imaging-plate diffractometer and a four-circle diffractometer, respectively (MoKα radiation, graphite monochromator).

Incorporation of rubidium cations (Rb<sup>+</sup>) with oxalic anions (C<sub>2</sub>O<sub>4</sub><sup>2-</sup>) establishes two-dimensional layer structure; water molecule (H<sub>2</sub>O)<sup>0</sup> intercalates into the layers. Conformation of thallium cations (Tl<sup>+</sup>) to seven oxygen atoms of oxalic anions sets up the two-dimensional layer structure, which is similar with that of rubidium oxalate. Except for water molecule, thallium oxalate is isotopic to rubidium oxalate (this study) and potassium oxalate K<sub>2</sub>(C<sub>2</sub>O<sub>4</sub>)·H<sub>2</sub>O[1].

Bond valence analysis of these compounds reveals that (H<sub>2</sub>O)<sup>0</sup> moderates the Lewis basicity of oxalic anion (C<sub>2</sub>O<sub>4</sub><sup>2-</sup>: 0.167). This moderated Lewis basicity matches the Lewis acidity of potassium and rubidium cation (K<sup>+</sup>: 0.126, Rb<sup>+</sup>: 0.124)[2], respectively: the valence-matching principle is satisfied.

[1] Sequeira A., Srikanta S., Chidambaram R. *Acta Cryst.*, 1970, **B26**, 77. [2] Brown I. D. *Acta Cryst.*, 1988, **B44**, 545.

**Keywords:** oxalate mineral, bond-valence, hydrogen bonds

**P.10.05.8***Acta Cryst.* (2005). A61, C377**Structure and Composition Properties of Mica from Ilmen Reserve Granite Pegmatites**

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Micas, picked up A) along the edges and B) from the center of five granite pegmatites were studied from Ilmen reserve (Ural, Russia), using the optical, X-ray, microprobe (SEM and TEM) and electron diffraction methods.

The 2V measurements showed two different types of mica: A) with homogeneity 2V distribution on all area of one crystal B) heterogeneous 2V distribution, which is also showed by X-ray and microprobing analyses. Some fluctuations in Ng' and Nm' significance are caused by heightened content of F, Fe and Li. слюды 2M<sub>1</sub>, 1M, 3T and M<sub>d</sub> polytypes were detected by X-ray powder diffraction. For all that optical homogeneity of micas were 2M<sub>1</sub> polytype, while the heterogeneous micas with intermediate *a* and *b* sense between di- and trioctahedral structures are presented in 1M and 3T polytypes. Fe-micas were an Md modification. Diagnostic and division of 1M and 3T modification in powder samples were made on intensity and position of 111 and 021 peaks.

Crystal formula was calculated on bulk chemical composition, SEM and TEM analyses, taking into account X-ray and optical data. Blocks with different composition possess so hard as smooth edges in one crystal, were detected for heterogeneous samples using the TEM and electron diffraction data. Interesting is, that two neighbor area have different ultimate mineral members. This complex of analyses allows the dioctahedral mica with heightened Fe content carry on muscovite and celadonite, trioctahedral mica – on flogopyte and annite and the Li-Fe mica on trillionite.

As a result the homogenous 2M<sub>1</sub> mica with utmost di- or trioctahedral composition on board and heterogeneous 1M and 3T polytypes with heightened content Li, F and Fe on center of pegmatites were exposed.

Prescribed specialities of homogeneity and heterogeneous structure and composition in mica can be used as typomorphic indicators of granite pegmatites crystallization conditions.

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**Keywords:** X-ray, mica, granite pegmatites

**P.10.05.9***Acta Cryst.* (2005). A61, C377-C378**XRD and HRTEM Structural Analyses of Antigorite Polysomes**

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The modulated crystal structure of two common antigorite polysomes, the "odd" *m*=17 and the "even" *m*=16 polysomes (*m* being the number of tetrahedral modules in a wave), have been recently refined by XRD using single crystals from the Val Malenco Italian Alps. The *m*=17 polysome [1] has *Pm* symmetry, cell constants *a*=43.505(6), *b*=9.251(1), *c*=7.263(1) Å, β=91.32(1)°, and consists of a wavy 1:1 layer, curled on the *b*-axis and modulated along [100]. Sixteen **O**-modules form a continuous **O**-sheet. Seventeen **T**-modules link the **O**-sheet at the concave sides, forming a "short" half-wave, and "long" half-wave with opposite polarities. The **T**-sheet shows 6-membered tetrahedral rings, like in lizardite, which reverse polarity via alternating 6-reversals (6-membered tetrahedral rings) and 8-reversals (8- and 4-membered tetrahedral rings, coupled along **b**).

The "even" *m*=16 polysome has *C*/2*m* symmetry, cell constants *a*=81.664(10), *b*=9.255(5), *c*=7.261(5) Å, β=91.41(5)°, and the same structural configuration of the *m*=17 polysome, but for the even number of **T**-modules, which makes the halfwaves symmetric, and the periodic **b**/2 shifting of the 8-reversal position, which makes the cell double and *C*-centered. First neighbors interactions are very similar in the two polysomes, and similar to lizardite [2].

HRTEM investigations carried on the same batch of crystals, confirm the presence of *P* and *C* structures, and show that the scheme

depicted above holds also for the  $m=15$  (P) and the  $m=18$  (C) polysomes.

[1] Capitani G.C., Mellini M., *Am. Mineral.*, 2004, **89**, 147. [2] Mellini M., Viti C., *Am. Mineral.*, 1994, **79**, 1194.

**Keywords:** antigorite, structure, polysomatism

#### P.10.05.10

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##### Two kinds of Superstructures of (Ge, Si) Wollastonite

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Two kinds of superstructures of wollastonite whose compositions were Ca(Ge<sub>0.55</sub>, Si<sub>0.45</sub>)O<sub>3</sub> and Ca(Ge<sub>0.65</sub>, Si<sub>0.35</sub>)O<sub>3</sub> were synthesized. The former is called 0.55Wo and the latter 0.65Wo in this abstract. The lattice constants of 0.55Wo and 0.65Wo are: (angstrom, degree)  $a=2X8.016(4)$ ,  $2X8.038(5)$ ,  $b=7.421(1)$ ,  $7.451(1)$ ,  $c=7.157(2)$ ,  $7.194(1)$ ,  $\alpha=90.08(2)$ ,  $89.93(1)$ ,  $\beta=94.86(2)$ ,  $94.85(2)$ ,  $\gamma=103.44(2)$ ,  $103.34(1)$ , respectively. Though both of them are twice as large as the basic structure of wollastonite, those structures are different. Basically, 0.55Wo has the two units of the basic-wollastonite along a-axis and has no other partial structures in it. The largest difference between the basic wollastonite and 0.55Wo is the fact that the basic wollastonite has an inversion center but 0.55Wo has not it. Therefore, the basic wollastonite includes only 15 independent atoms but 0.55Wo includes the 60 atoms in the unit cell. The final R-value was 6% after the structure refinement of 0.55Wo. 0.65Wo has the two units of the basic-wollastonite along a-axis, too. However, the X-ray diffraction pattern of 0.65Wo was largely different from that of 0.55Wo and it showed the extinction rule of the pseudo-C lattice. The structure model having the stacking of the basic-wollastonite unit with  $b/2$  displacement were tried and the R-value was 9% after the structure refinement.

**Keywords:** structure analysis, superstructure, wollastonite

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##### Light-Induced Degradation Dynamics in Realgar (As<sub>4</sub>S<sub>4</sub>)

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Light-induced degradation in realgar (As<sub>4</sub>S<sub>4</sub>) has been studied by means of four-circle XRD and XPS. Because of the alteration of realgar exposed to light, the  $a$  lattice parameter and  $c \sin\beta$  value increase linearly, which leads to a continuous increase of the unit cell volume. Nevertheless, no correlation exists between the continuous increase of the unit cell volume and the bond distance variations in As<sub>4</sub>S<sub>4</sub> molecules. The most pronounced change was observed in the distance between centroids in As<sub>4</sub>S<sub>4</sub> cage. The spread of As<sub>4</sub>S<sub>4</sub> intermolecular distances increases continuously from 5.642 Å to 5.665 Å, which directly affects the unit cell volume expansion of realgar. In addition, the O1s peak increases rapidly after light exposure. The result substantiates the reaction proposed by Bindi et al. [1]. That is, realgar is transformed into pararealgar if oxygen exists and produces As<sub>4</sub>S<sub>5</sub> molecule. The additional S atom at As<sub>4</sub>S<sub>4</sub> molecule contributes to anisotropic expansion for the  $a$  and  $c$  axes because the direction for the additional S atom points toward [4-14] in the unit cell. Furthermore, an S atom in the As<sub>4</sub>S<sub>5</sub> molecule is released from one of equivalent As-S-As linkages in As<sub>4</sub>S<sub>5</sub>; it turns itself into the As<sub>4</sub>S<sub>4</sub> molecular of pararealgar. The free S atom is re-attached to another As<sub>4</sub>S<sub>4</sub> (realgar type) molecule, and reproduces an As<sub>4</sub>S<sub>5</sub> molecule. The reproduced As<sub>4</sub>S<sub>5</sub> molecule turns itself into the As<sub>4</sub>S<sub>4</sub> molecular of pararealgar again. This cycle that realgar is indirectly transformed into pararealgar via As<sub>4</sub>S<sub>5</sub> molecule is repeated during light exposure.

[1] Bindi L., Popova V., Bonazzi P., *Can. Mineral.*, 2003, **41**, 1463.

**Keywords:** single-crystal XRD, mineral physics, photochemistry

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##### A Revision of the Structure and Crystal-Chemistry of the Arrojadite-Group

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Arrojadite, dickinsonite and sigismundite define a group of phosphates with a very complex structure and crystal-chemistry, which has not been satisfactorily characterized so far. Previous works described the structure in the  $C2/c$  space group which implies the presence of half-occupied P1 phosphate groups sharing basal faces and with the apical oxygen atoms pointing in opposite directions, and of partial occupancy at 10 out of the 49 sites in the asymmetric unit. Electron microprobe analyses suffer from severe interference problems due to the coexistence of P, Fe and F.

Single-crystal refinement, Raman spectroscopy and well-controlled *in situ* microchemical analysis of selected samples allowed a better interpretation of the structure and crystal-chemistry of the arrojadite group. The correct space group is  $Cc$ , and the lowering in symmetry implies recognition of strong cation ordering among the 21 cationic sites besides the 12 PO<sub>4</sub> groups. The correct number of OH groups  $pfu$  is 3 and not 2 (2 OH are bonded to octahedral sites and 1 to a PO<sub>4</sub> group), and a selectivity is shown for F incorporation at one of these sites. The active exchange vectors were identified, and a new nomenclature has been proposed on crystal-chemical basis.

The  $C2/c$  derives from the presence in some samples of  $Cc$  domains related by a diad along  $b$  plus a  $1/2$  shift along  $c$ . This feature cannot be treated by X-ray structure refinement, and may be related to growth defects occurring during high- $T$  crystallisation (in pegmatites).

**Keywords:** phosphate crystal-chemistry, mineralogy, crystal-structure

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##### Supramolecular Arrangement of Tensioactive Molecules in Interlayer Spaces of Two Organophilic Clays

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Studies of mixed systems using clays and tensioactive molecules (TA) are of particular importance in many fields such as medicine, pharmacy and cosmetics. The substitution of the initial cations and intercalation of cationic organic molecules in the interlayer spaces of the clay structure lattice provide new physicochemical properties. The aim of the present study consist in a structural characterisation of two different purified and organophilic clays, mainly made of smectite with a marked montmorillonitic nature, as a function of the inserted TA quantity using high-resolution X-ray powder diffraction. Two different TA are used, the benzyltétradécyltriméthylammonium (BTDMA) and the benzyldodécyltriméthylammonium (BDDMA).

**Keywords:** organomineral materials, powder X-ray diffraction, supramolecular assemblies

#### P.10.05.14

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##### Ti-investigation by XPS of Biotite from Metapelites of El Joyazo (Spain)

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